

# Nutrients dynamics in the main river basins of the centre-southern region of Chile

Jaime Pizarro<sup>a,\*</sup>, Pablo M. Vergara<sup>a</sup>, José A. Rodríguez<sup>a</sup>, Pedro A. Sanhueza<sup>a</sup>, Sergio A. Castro<sup>b</sup>

<sup>a</sup> Departamento de Ingeniería Geográfica, Facultad de Ingeniería, Universidad de Santiago de Chile, Av. Lib. B. O'Higgins 3363, Santiago, Chile

<sup>b</sup> Departamento de Biología, Facultad de Química y Biología, Universidad de Santiago de Chile, Av. Lib. B. O'Higgins 3363, Santiago, Chile

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## ABSTRACT

Chilean basins have long been exposed to nutrient discharges from human activities and land use changes. A historical seasonal  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P database of the last 23 years of the main nine rivers of central-southern region of Chile was analysed. Generalized additive models indicated that annual trends in  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P are nonlinear. River basins such as Bío-Bío, Bueno, Imperial, Maule, Rapel and Valdivia showed a clear increase in  $\text{NO}_3^-$ -N, while  $\text{PO}_4^{3-}$ -P increased only in the Rapel and Maule basins. Although no seasonal difference in  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P was found in the analysed basins, there was a negative relation between these nutrients and water flow. Sampling stations with high  $\text{NO}_3^-$ -N concentration were found mostly in sub-basins located in the "central valley" of central Chile, while several  $\text{PO}_4^{3-}$ -P "over-concentrated" sampling stations were located mostly upstream. If  $\text{NO}_3^-$ -N emissions into Chilean river basins continue at current rates it is probable that the concentration of this nutrient will tend to match that of the most "polluted" rivers around the world.

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## 1. Introduction

Phosphorous and nitrogen are two essential nutrients for the aquatic biota, and limits their primary productivity [1]. However, overload of these nutrients accelerates the eutrophication process, and deteriorates river water quality [1].

In the last century, wastes from anthropogenic activities (e.g., domestic wastes, agriculture) have strongly affected most freshwater ecosystems [2–9]. As a result, the total river-load of dissolved nitrate–nitrogen ( $\text{NO}_3^-$ -N) has increased nearly six times and the total phosphate–phosphorous ( $\text{PO}_4^{3-}$ -P) nine times over pre-industrial levels in the main rivers around the world [10,11]. The latter has motivated researchers to assess the impacts of anthropogenic activities on the deterioration of water quality in local, regional and continental river systems using image processing techniques and extensive monitoring programs [9,12–15].

The central-southern section of Chile includes several latitudinal distributed river basins that originate at the top of the Andean mountains flowing into the Pacific Ocean (Fig. 1). Historically, the main productive activities in Chile's central-southern region have been agriculture, forestry and livestock (Table 1), whose intensities have been increasing over time as population growth in this region. Accordingly, we should expect a significant increase in the levels of nitrate and phosphate in Chilean rivers. Previous studies have

shown that the concentration and fluxes of dissolved nitrogen in small watersheds of Chile's central-southern region is significant increased by agriculture, forestry and livestock [16–18]. Regardless of the contribution of the later findings to the understanding of the anthropogenic impact upon nutrient, it is necessary to know the long-term temporal behaviour of nutrients in the main river basins of Chile. In this study we analyzed an historical seasonal  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P database of the nine main rivers in Chile's central-southern region over the last 23 years (Table 1). Our interest was three-fold: first we evaluated the annual and seasonal dynamics in concentration of both nitrate–nitrogen ( $\text{NO}_3^-$ -N) and phosphate–phosphorous ( $\text{PO}_4^{3-}$ -P); second we identified rivers where  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P concentration has increased in the study period by comparing their annual and seasonal trends in these rivers; and finally, we proposed possible mechanisms that explain these trends.

## 2. Materials and methods

### 2.1. Study site

We studied the nine main basins of central-southern Chile (from  $33^\circ 53'$  to  $39^\circ 50'$  lat S., Fig. 1). The climate of these basins varies latitudinally, from a semi-desertic climate in the north to a temperate rain in the south (with mean annual precipitation ranging from 300 mm to 1800 mm). As a consequence of this climate gradient, water flows tend to be larger in the southern river basins. Furthermore, the dry season occurs in late spring and summer and the wet season in late autumn and winter. According to bio-

\* Corresponding author. Tel.: +56 2 7182226.

E-mail address: [jaime.pizarro@usach.cl](mailto:jaime.pizarro@usach.cl) (J. Pizarro).

**Table 1**

Summary of studied basins, including their area (km<sup>2</sup>), percentage (%) of landscape area used for agriculture, forestry or livestock, as well as the mean annual water flow (m<sup>3</sup> s<sup>-1</sup>) during the study period.

Basin	Area	Agriculture	Forestry	Livestock	Water flow
Rapel	13,695	0.22	0.01	0.05	30.8
Mataquito	6,190	0.17	0.03	0.06	46.5
Maule	20,295	0.22	0.09	0.04	78.1
Itata	11,294	0.01	0.07	0.01	67.8
Bío-Bío	24,264	0.03	0.2	0.07	154.0
Imperial	12,763	0.01	0.12	0.07	55.6
Toltén	8,398	0	0.04	0.3	114.0
Valdivia	10,275	0	0.13	0.28	145.4
Bueno	15,637	0	0.03	0.44	99.3

climate conditions, these river basins have historically differed in their human activities, with the lands of the northern basins used mostly for agriculture and the southern basins for livestock (Table 1).

2.2. Sampling and analysis

Sampling data belongs to the National River Monitoring Network of the Dirección General de Aguas (DGA) of the Ministry of Public Works of Chile. The DGA's sampling stations are systematically distributed along the main tributaries of each basin (sub-basins), covering from the river head to the river mouth (Fig. 1, Table 2). Water samples were collected on a seasonal basis using high density polyethylene (HDPE) containers (500 mL). NO<sub>3</sub><sup>-</sup>-N was measured according to the methodology described by [19], using molecular absorption spectroscopy with a detection limit of 0.002 mg L<sup>-1</sup>. PO<sub>4</sub><sup>3-</sup>-P was measured by method 4500 as suggested by SMEWW [20] using molecular absorption spectroscopy with a detection limit of 0.003 mg L<sup>-1</sup>.

**Table 2**

Number of sampling stations, data records of NO<sub>3</sub><sup>-</sup>-N and PO<sub>4</sub><sup>3-</sup>-P and sampling years for the nine river basins analyzed in this study.

Basin	NO <sub>3</sub> <sup>-</sup> -N			PO <sub>4</sub> <sup>3-</sup> -P		
	Sampling stations (n)	Data records (n)	Years*	Sampling stations (n)	Data records (n)	Years*
Rapel	18	764	20	18	741	21
Mataquito	9	307	19	10	324	20
Maule	20	852	19	21	893	21
Itata	16	681	21	16	694	22
Bío-Bío	14	755	21	14	837	22
Imperial	12	779	20	12	851	22
Toltén	7	413	20	7	447	22
Valdivia	16	680	21	16	714	23
Bueno	20	822	21	20	863	22

\*Number of sampling years.

2.3. Statistical analysis

To assess yearly and seasonal trends in NO<sub>3</sub><sup>-</sup>-N and PO<sub>4</sub><sup>3-</sup>-P we used Generalized Additive Mixed Models (GAMM), which allowed us: (1) to analyze nonlinear temporal trends, fitting smooth curves (penalized cubic regression splines); (2) to analyze hierarchical nested data such as those in this study [21,22]; Random errors associated with sampling stations were spatially nested within sub-basins, and those associated with sub-basins within river basins. In addition, we included random intercepts for year and season to correct for non-independence of observations within the same stations [22]. Previously, we normalized NO<sub>3</sub><sup>-</sup>-N and PO<sub>4</sub><sup>3-</sup>-P data by log transforming.

In order to analyse annual trends, and considering that NO<sub>3</sub><sup>-</sup>-N and PO<sub>4</sub><sup>3-</sup>-P could be affected by other independent variables (such as water flow and station), we developed models with different variable combinations, including water flow as a continuous

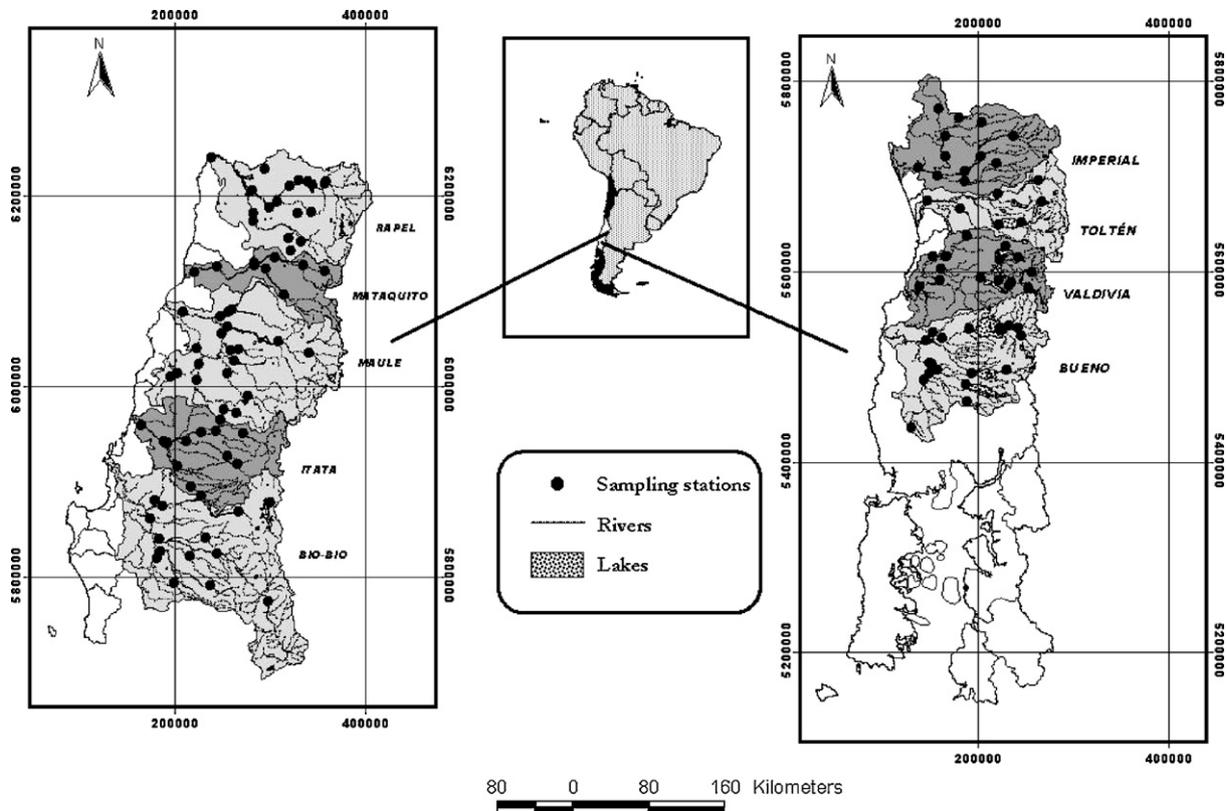


Fig. 1. Map of the nine river basins studied in the centre-southern Chilean region.

**Table 3**

Annual mean  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P concentrations and their standard errors (SE) in the largest nine river basins of central-southern Chile. For each river basin the percentage of sampling stations with significant level concentration (% significant, see Section 3) as well as the mean elevation (in m.a.s.l.) of all sampling stations (S elevation) and elevation of significant sampling stations (SS elevation) are shown.

Basin	Mean	SE	% significant	S elevation	SS elevation
<b><math>\text{NO}_3^-</math>-N</b>					
Rapel	1.57	0.31	83.3	242.8	206.7
Mataquito	0.55	0.22	50.0	334.5	29.0
Maule	0.49	0.22	71.4	151.9	122.7
Itata	0.35	0.09	25.0	144.8	49.5
Bío-Bío	0.21	0.12	21.4	171.9	46.0
Imperial	0.4	0.16	58.3	155.9	107.7
Toltén	0.11	0.08	0.0	237.1	–
Valdivia	0.31	0.11	12.5	114.3	0.0
Bueno	0.18	0.11	20	175.5	11.3
<b><math>\text{PO}_4^{3-}</math>-P</b>					
Rapel	0.23	0.16	27.8	242.8	253.6
Mataquito	0.29	0.18	44.4	338.3	36.3
Maule	0.24	0.17	50.0	158.5	129.0
Itata	0.41	0.14	81.3	144.8	174.4
Bío-Bío	0.15	0.1	21.4	171.9	26.7
Imperial	0.23	0.1	33.3	155.9	212.5
Toltén	0.23	0.11	57.1	237.1	390.0
Valdivia	0.10	0.07	12.5	114.3	0.0
Bueno	0.12	0.08	10.0	175.5	0.0

**Table 4**

General additive mixed models for  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P ranked according to their Akaike Information Criterion (AIC). Adjusted  $R^2$  is also shown for each model. Best parsimonious models are indicated by bold cases.

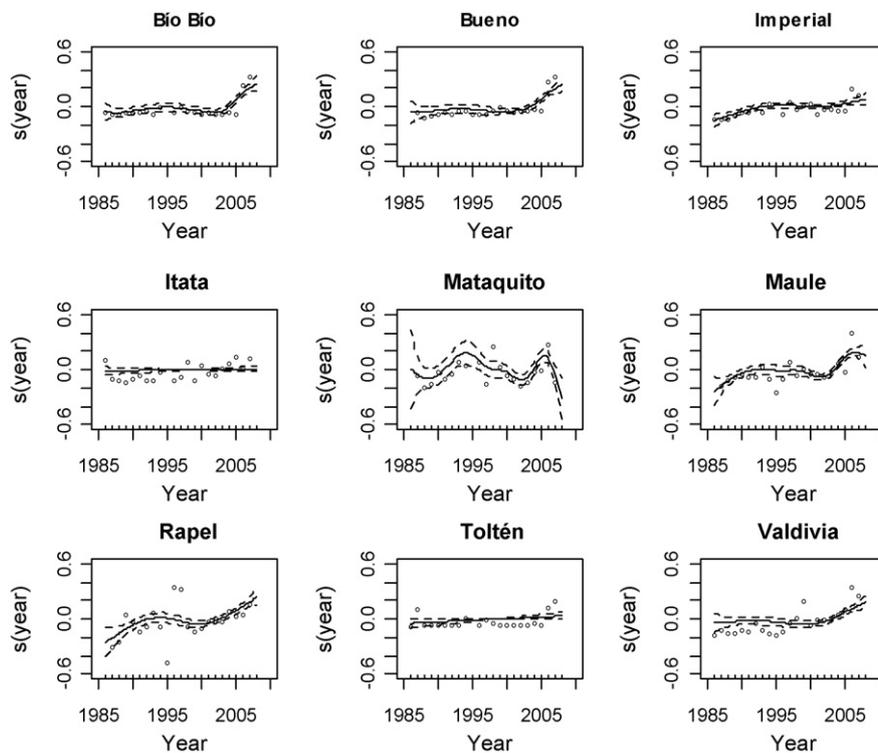
Model variables*	AIC	$\Delta$ AIC	$R^2$
<b><math>\text{NO}_3^-</math>-N</b>			
$s(\text{years}) + \text{flow} + \text{season} + \text{sampling station}$	1298.8	<b>0.0</b>	0.59
$s(\text{years}) + \text{season} + \text{sampling station}$	1302.2	<b>-3.4</b>	0.59
$s(\text{years}) + \text{flow} + \text{sampling station}$	1305.6	-6.8	0.59
$s(\text{years}) + \text{sampling station}$	1306.4	-7.6	0.59
$s(\text{years}) + \text{flow} + \text{season} + \text{river basin}$	1652.4	-353.6	0.22
$s(\text{years}) + \text{flow} + \text{river basin}$	1654.7	-355.9	0.22
$s(\text{years}) + \text{season} + \text{river basin}$	1655.1	-356.3	0.22
$s(\text{years}) + \text{river basin}$	1655.3	-356.6	0.22
$s(\text{years}) + \text{flow} + \text{season}$	1659.0	-360.2	0.04
$s(\text{years}) + \text{flow}$	1661.3	-362.5	0.04
$s(\text{years}) + \text{season}$	1661.8	-363.0	0.04
<b><math>\text{PO}_4^{3-}</math>-P</b>			
$s(\text{years}) + \text{flow} + \text{season} + \text{sampling station}$	1028.6	<b>0.0</b>	0.28
$s(\text{years}) + \text{flow} + \text{sampling station}$	1036.3	-7.8	0.28
$s(\text{years}) + \text{season} + \text{sampling station}$	1040.9	-12.3	0.28
$s(\text{years}) + \text{sampling station}$	1049.1	-20.5	0.28
$s(\text{years}) + \text{flow} + \text{season} + \text{river basin}$	1071.9	-43.3	0.25
$s(\text{years}) + \text{flow} + \text{season}$	1078.6	-50.0	0.24
$s(\text{years}) + \text{flow} + \text{river basin}$	1080.1	-51.5	0.25
$s(\text{years}) + \text{flow}$	1086.3	-57.8	0.23
$s(\text{years}) + \text{season}$	1090.1	-61.5	0.23
$s(\text{years}) + \text{river basin}$	1091.7	-63.1	0.25

\*  $s(\text{years})$  represent a smooth function of time.

covariate and season and sampling station as factors with multiple levels (including each season and each sampling station, respectively). We used the Akaike's Information Criterion to rank and select the most parsimonious models. The Akaike Information Criterion (AIC) is a measure of the model deviance penalized by its number of parameters so that the preferred model is the one with the lowest AIC value [23]. We considered that a model is parsimonious if its  $\Delta\text{AIC} \leq 4$ . Water flow was standardized at the basin level to control the differences in water flow between basins.

### 3. Results

$\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P concentration varied differently among basins, with a mean  $\text{NO}_3^-$ -N concentration per basin ranging from 0.11  $\text{mg L}^{-1}$  to 1.57  $\text{mg L}^{-1}$ , corresponding to Toltén and Rapel basins, respectively (Table 3). The mean concentration of  $\text{PO}_4^{3-}$ -P per basin ranged from 0.10  $\text{mg L}^{-1}$  to 0.41  $\text{mg L}^{-1}$ , corresponding to the Valdivia and Itata basins, respectively (Table 3).



**Fig. 2.** Annual concentration ( $\text{mg L}^{-1}$ ) trends of  $\text{NO}_3^-$ -N in the nine river basins studied in the centre-southern Chilean region. Y-axis represents log-transformed partial residuals, with smoothing GAMM curves  $[s(\text{year})]$  represented by continuous lines while upper and lower confidence intervals by dashed lines. Circles represent the observed log-transformed mean in annual concentration of  $\text{NO}_3^-$ -N per basin.

**Table 5**

Estimated degrees of freedom (edf) and significance ( $p$ -value) of smooth functions of time (year) for the best models predicting  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P concentration in river basins of central-southern Chile.

Basin	$\text{NO}_3^-$ -N		$\text{PO}_4^{3-}$ -P	
	edf	$p$ -value	edf	$p$ -value
Rapel	4.3	<0.001	6.6	<0.001
Mataquito	7.3	<0.001	8.1	<0.001
Maule	6.5	<0.001	8.0	<0.001
Itata	1.0	0.793	8.4	<0.001
Bío-Bío	5.6	<0.001	5.5	<0.001
Imperial	3.2	<0.001	7.2	<0.001
Toltén	1.0	0.136	5.1	<0.001
Valdivia	3.7	<0.001	6.5	<0.001
Bueno	4.5	<0.001	7.1	<0.001

### 3.1. Nitrate trends

According to AIC values, the best two models explaining  $\text{NO}_3^-$ -N included the effects of year, water flow, season and sampling station (Table 4). These best two models accounted for 60% variance in  $\text{NO}_3^-$ -N. For all basins, annual trends in  $\text{NO}_3^-$ -N were mostly nonlinear, as shown by smooth functions fitted to the data (Fig. 2, Table 5). From these plots, Bío-Bío, Bueno, Imperial, Maule, Rapel and Valdivia showed an evident increase in  $\text{NO}_3^-$ -N during the sampling period (Fig. 2). For these basins,  $\text{NO}_3^-$ -N did not increase at a constant rate, since the increase rate was greater after 2000 (Bío-Bío, Bueno, Valdivia, Maule, Rapel and Imperial increased 9.89, 7.56, 4.32, 2.26, 1.85, 1.43 times, respectively; Fig. 2). For the Itata and Tolten basins there was no clear temporal trend in  $\text{NO}_3^-$ -N, since their smooth functions were not significant (Table 5). The Mataquito basin, however, shows a clear increase between 1987 and 1995, and during 2006. During these periods the mean  $\text{NO}_3^-$ -N value was 1.25 times and 2.50 times larger,

**Table 6**

Regression coefficients (estimate) with their standard errors (SE),  $t$ -values ( $t$ ), and  $p$ -values for independent variables of the best models predicting  $\text{NO}_3^-$ -N and  $\text{PO}_4^{3-}$ -P concentration in river basins of central-southern Chile (see Table 4).

Variable	Estimate	SE	$t$	$p$ -value
$\text{NO}_3^-$ -N				
Flow	-0.01	0.004	-2.3	0.021
Autumn	0.013	0.011	1.2	0.235
Spring	-0.014	0.011	-1.3	0.186
Summer	-0.024	0.013	-1.8	0.066
$\text{PO}_4^{3-}$ -P				
Flow	-0.017	0.005	-3.8	<0.001
Autumn	-0.004	0.01	-0.4	0.671
Spring	0.016	0.01	1.6	0.114
Summer	-0.021	0.013	-1.6	0.109

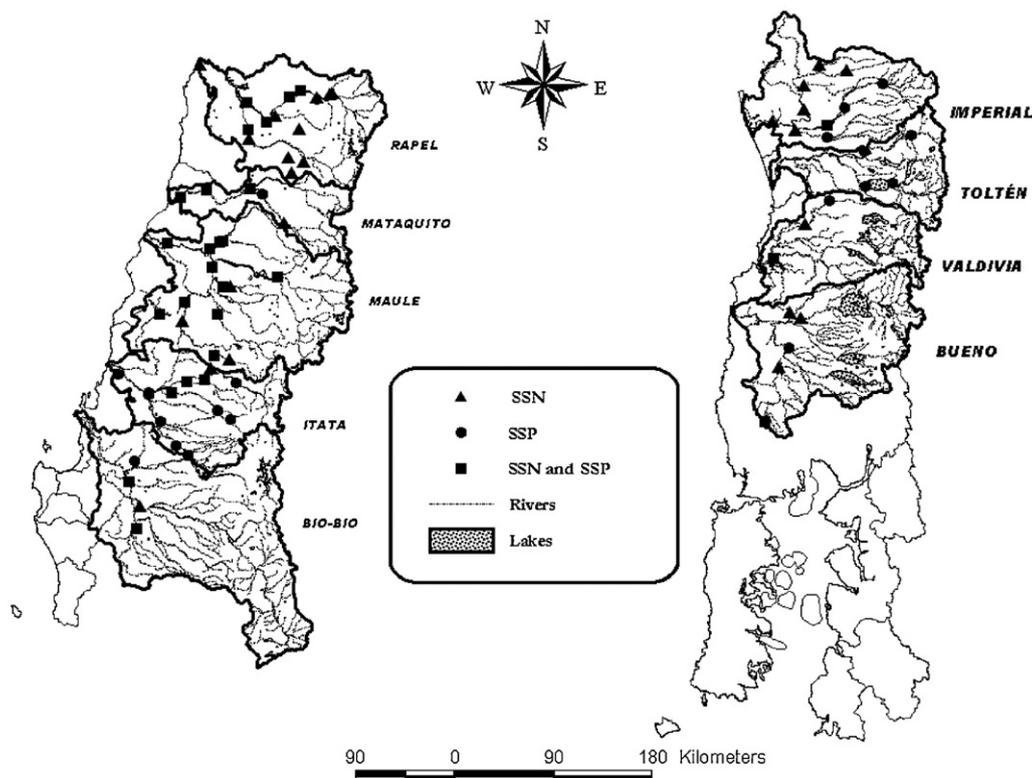
\*  $p < 0.05$ , \*\*\*  $p < 0.001$ .

respectively, than the mean value during the total sampled period (Fig. 2).

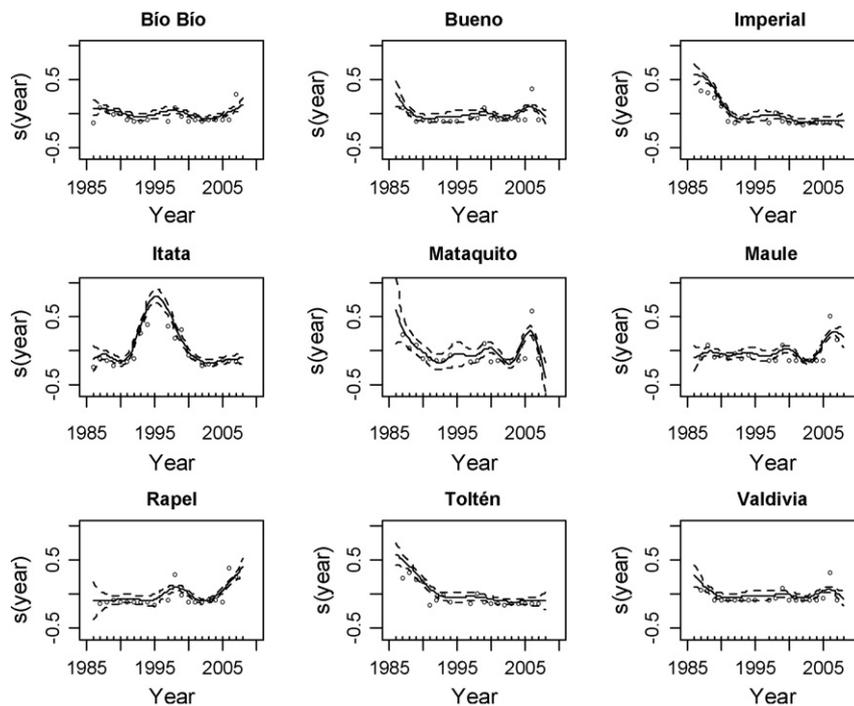
Model results indicated that there is no seasonal difference in  $\text{NO}_3^-$ -N in the analysed basins (Table 6). We found a negative relation between  $\text{NO}_3^-$ -N and water flow (Table 6). Furthermore, in 56 sampling stations (41.0%)  $\text{NO}_3^-$ -N concentration was significantly higher than in the reference station which was that with the lowest annual mean  $\text{NO}_3^-$ -N ( $0.025 \text{ mg L}^{-1}$ ). The Rapel and Maule basins had the largest proportion of these “over-concentrated” sampling stations (Table 3; Fig. 3). Moreover, all the basins over-concentrated sampling stations were located mostly in the lower section of all the studied basins (Table 3; Fig. 3).

### 3.2. Phosphate trends

The best model accounting for  $\text{PO}_4^{3-}$ -P included the effect of year, water flow, season, and sampling station, which only explained about 30% variance in  $\text{PO}_4^{3-}$ -P (Table 4). For all basins,



**Fig. 3.** Map of the nine river basins studied in the centre-southern Chilean region showing the sampling stations with significant level concentration of  $\text{NO}_3^-$ -N (SSN) and  $\text{PO}_4^{3-}$ -P (SSP).



**Fig. 4.** Annual concentration ( $\text{mg L}^{-1}$ ) trends of  $\text{PO}_4^{3-}\text{-P}$  in the nine river basins studied in the centre-southern Chilean region. Y-axis represents log-transformed partial residuals, with smoothing GAMM curves [ $s(\text{year})$ ] represented by continuous lines while upper and lower confidence intervals by dashed lines. Circles represent the observed log-transformed mean in annual concentration of  $\text{PO}_4^{3-}\text{-P}$  per basin.

annual trends in  $\text{PO}_4^{3-}\text{-P}$  were mostly nonlinear and significant as shown by smooth functions fitted to the data (Fig. 4, Table 5). From these plots, only Rapel and Maule showed an evident increase in  $\text{PO}_4^{3-}\text{-P}$  during the last years (after 2000) of the sampling period (Rapel and Maule increased 10.37 and 7.85 times, respectively, Fig. 4). The Itata basin, however, shows a clear increase between 1990 and 2000, a period in which the mean concentration was 1.9 times greater than the mean value during the total period (Fig. 4). Three basins, Imperial, Mataquito, and Tolten, showed a clear negative annual trend in  $\text{PO}_4^{3-}\text{-P}$  concentration during the sampling period (Fig. 4). In the Imperial and Tolten basins this decrease in  $\text{PO}_4^{3-}\text{-P}$  occurred mainly during the first 80 years of the sampling period (1986 and 1993, Fig. 4). The Bío-Bío, Bueno, and Valdivia basins did not show a clear trend even though their smooth functions were significant, (Fig. 4, Table 5).

Model results indicated that there is no seasonal difference in  $\text{PO}_4^{3-}\text{-P}$  in the analysed basins (Table 6). We found a negative relation between  $\text{PO}_4^{3-}\text{-P}$  and water flow (Table 6). Furthermore, in 47 sampling stations (35.6%)  $\text{PO}_4^{3-}\text{-P}$  concentration was significantly higher than the reference sampling station with the lowest annual mean  $\text{PO}_4^{3-}\text{-P}$  concentration ( $0.008 \text{ mg L}^{-1}$ ). The Itata basin had the highest proportion of these “over-concentrated” sampling stations (Table 3; Fig. 3). Moreover, in the Mataquito, Maule, Bío-Bío, Valdivia, and Bueno basins, the sampling stations with large  $\text{PO}_4^{3-}\text{-P}$  concentration were located mostly in the lower sections (Table 3; Fig. 3). In contrast, in the Itata, Rapel, Imperial, and Tolten basins the sampling stations with large  $\text{PO}_4^{3-}\text{-P}$  concentration were located mostly in the upper sections (Table 3; Fig. 3).

#### 4. Discussion

Contamination by  $\text{NO}_3^- \text{-N}$  and  $\text{PO}_4^{3-}\text{-P}$  is a global phenomenon affecting freshwater systems resulting from intensive human activities [24]. Chilean basins in the central-southern region have been largely affected by human activities over the last cen-

turies. Human population has increased dramatically over the last 30 years and land use changed drastically from small to large scale agriculture and forestry. Accordingly, our results indicate that  $\text{PO}_4^{3-}\text{-P}$  and  $\text{NO}_3^- \text{-N}$  inputs have increased during this time period at least two and six basins, respectively. Although the annual average  $\text{NO}_3^- \text{-N}$  concentration was higher in Rapel, while its most pronounced increases occurred in those basins that were less “polluted” at the beginning of the sampling period (e.g., Bío-Bío, Bueno and Valdivia), suggesting that N concentration in these basins may become important in the middle term. High  $\text{NO}_3^- \text{-N}$  concentration in the Rapel basin probably results from the large population density, tourism in the Rapel reservoir, and particularly the intensive industrial–agricultural activities (Fig. 3) [25,26]. The significant increase in the last 10 years in the Bueno and Valdivia basins can be explained by the fact that both basins share similar human activities (e.g., livestock), and in the last decade human population has increased greatly. Similarly, the Bío-Bío basin has experienced a large expansion in forestry and manufacturing industries in this time period.

Our results show that  $\text{PO}_4^{3-}\text{-P}$  concentration follows a nonlinear trend in most studied basins, but only Rapel and Maule showed a positive increase in the last decade. On the contrary, in several basins, such as Itata, Bío-Bío, Bueno, and Valdivia, we found variable fluctuations in  $\text{PO}_4^{3-}\text{-P}$  concentration, but we failed in finding any trend. Probably, the large increase in  $\text{PO}_4^{3-}\text{-P}$  between 1990 and 2000 in the Itata basin was due to the expansion of pine forest plantations during the 1990s decade. It should be noticed, however, that the best models for  $\text{PO}_4^{3-}\text{-P}$  concentration accounted for less than 30% of its variance as shown by the adjusted  $R^2$  (see Table 4). This means that variables independent of time, or not associated with water flow or sampling station, could be more important in predicting  $\text{PO}_4^{3-}\text{-P}$  concentration in Chile’s basins.

Models indicated that variance in  $\text{NO}_3^- \text{-N}$  was accounted for differences between sampling stations rather than by differences between basins (Table 4). On the contrary, the addition of the sam-

pling station as a factor did not improve the fit of  $\text{PO}_4^{3-}\text{-P}$  models. This implies that source emissions of  $\text{NO}_3^-\text{-N}$  are not homogeneously distributed along the basin, and therefore the efforts for controlling water quality should be focused only on those sub-basins in which  $\text{NO}_3^-\text{-N}$  concentration has increased significantly. Moreover, we found that sampling stations with a large concentration of  $\text{NO}_3^-\text{-N}$  were located in the lowland areas within each basin. Specifically, these sampling stations were distributed over the “central valley” of central-southern Chile. The central valley is a flat area where historically human activities such as industries, livestock, and urban areas have developed more. Nevertheless, activities within the central valley are not homogeneously distributed among the basins because of the latitudinal bio-climatic gradient and the fact that the northernmost basins, such as Rapel or Maule, have been exposed to human perturbation for a longer time.

According to our results,  $\text{PO}_4^{3-}\text{-P}$  and  $\text{NO}_3^-\text{-N}$  concentrations have different temporal dynamics among and within river basins. At the basin scale, there are more river basins exhibiting increasing trends  $\text{NO}_3^-\text{-N}$  than in  $\text{PO}_4^{3-}\text{-P}$  concentration. We found an important contribution of the sampling station in accounting for variation in  $\text{NO}_3^-\text{-N}$  concentration, but the sampling station was not an important factor for the  $\text{PO}_4^{3-}\text{-P}$  models (Table 4). Moreover,  $\text{NO}_3^-\text{-N}$  “over-concentrated” sampling stations were located mostly in the lower sections or sub-basins, while several  $\text{PO}_4^{3-}\text{-P}$  “over-concentrated” sampling stations were located upstream rather than in the lower basin sections (Table 3).  $\text{PO}_4^{3-}\text{-P}$  emissions may probably decrease downstream as a result of either the precipitation in phosphate compounds or photosynthetic processes.

We did not find a seasonal variation in  $\text{PO}_4^{3-}\text{-P}$  and  $\text{NO}_3^-\text{-N}$  concentration, contrasting with other studies carried out in other important river basins [8,27,28]. Seasonal differences of these parameters may be explained by soil loss and dilution in wet or flood seasons [9]. Water flow affected  $\text{PO}_4^{3-}\text{-P}$  and  $\text{NO}_3^-\text{-N}$  concentration negatively. Therefore, since water flow in Chilean rivers studied are dominated by two processes, snow melting in spring–summer and rain in winter, our results suggest that dilution could be the most important water flow-associated mechanism explaining the concentration of  $\text{PO}_4^{3-}\text{-P}$  and  $\text{NO}_3^-\text{-N}$ . We must also admit that total nitrogen may be present as  $\text{NH}_4^+\text{-N}$ , especially in basins with a large number of people and more agricultural activities, such as basins located further north of the Bío-Bío basin (see Fig. 3). However, it is necessary to highlight that nitrate, as quantified in this study, represents the most important contributor to total dissolved nitrogen in the world’s largest rivers (nearly 75%) [29].

The water quality of the river basins of central-southern Chile, in terms of their  $\text{NO}_3^-\text{-N}$  concentration, ranges between  $0.1\text{ mg L}^{-1}$  to  $1.6\text{ mg L}^{-1}$ , which is relatively similar to the concentration ranges of the world’s largest rivers and several watershed of the USA [29]. Of these rivers, the northernmost Rapel, Mataquito, and Maule basins show the highest concentrations ( $>0.5\text{ mg L}^{-1}$ ) and are comparable to European basins [10,29] or Chinese rivers [8]. Unfortunately, our results are not comparable with other studies carried out in small Chilean watersheds [16–18], which could have increased our understanding about how local anthropogenic activities (e.g., logging or fire) contribute in explaining the current nutrient levels in the main rivers of southern Chile. However, our results showed that if  $\text{NO}_3^-\text{-N}$  emissions into Chilean river basins continue to increase at current rates, it is probable that its concentration will tend to match that of the most “polluted” rivers around the world. Therefore, we propose that our results can be used as a reference source for future monitoring programs, ecosystem research, and mitigation projects in river basins of central-southern Chile.

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